

**ANCIENT WATER MANAGEMENT AND THE EVOLUTION OF THE LATE HOLOCENE WETLANDS. FIRST PALEOECOLOGICAL EVIDENCE FROM PREHISPANIC RAISED FIELDS OF URABÁ, NORTHWESTERN SOUTH AMERICA.**

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## **Summary**

The raised fields discovered recently in the Gulf of Urabá, northwestern Colombia, extends more than 135,000 ha among the floodplains of the rivers León, Suriquí and Tumaradocito in the Chocó biogeographic region. To understand the mechanisms by which people use the wetlands under climate change of the Late Holocene, the paleoenvironmental and cultural conditions were studied using artifacts, soil micromorphology, geochemical, chronostratigraphic and palynological analyses at El Vergel archaeological site. The aim of this study is to discuss the origin of raised fields development in the León river floodplain and its relation with progressive drought, groundwater and wetland management for permanent human occupation. The results suggest that the raised fields were built around the IX century CE, during a period marked by decreased precipitation, probably related with Medieval Warm Period. The hydrogeology reveals two aquifers in the region with some shallow springs in the floodplain where earthworks are located. The poor stratigraphic demarcation of the ridges and well dug channels, with no evidence for agriculture whatsoever, support the idea that some raised fields was strategic for distributing groundwater and rainwater over a large area and thus preserving the productivity of the wetland for fishing and hunting.

1 **Key words:** Wetlands, Raised fields, Uraba, Prehispanic cultures, Geoarchaeology,  
2 Holocene, paleoecology.

3

#### 4 **1. Introduction**

5 The massive occurrence of prehispanic raised fields near the Colombian-Panamanian border  
6 in the Gulf of Urabá, comprise an extensive wetland across the middle basin of the León  
7 River with pervasive earthworks of prehispanic cultures (Posada et al., 2019). The regional  
8 occurrence of these raised fields indicate a long term human interactions with wetland  
9 ecosystems and could support the statements about intercultural relationships among peoples  
10 from the Isthmus, northwestern Andes and Colombian Caribbean (Bray, 1990; Castillo,  
11 1988; Cooke and Sánchez, 2001; Fonseca and Cooke, 1993; Helms, 1979; Martin and  
12 Sánchez, 2007; Mendizábal et al., 2021; Piazzini, 2020; Reichel - Dolmatoff and Dussán de  
13 Reichel, 1961).

14

15 The raised fields are an ancient technology widespread in Central and South America for  
16 water management, usually for improving agricultural production (Comptour et al., 2018;  
17 Denevan, 2006, 1970). However, the regional applications of this technology vary with the  
18 environment, culture and the ecosystem services that peoples aimed for preservation.  
19 Previous works have shown the particular paleoenvironmental and cultural conditions that  
20 influenced fluvial ecosystems and productive behavior of societies around the widespread  
21 use of raised fields in South American floodplains (Iriarte et al., 2010; Ordóñez, 2006;  
22 Rodrigues et al., 2018; Rostain, 2010). These works highlight the diversity of functions and  
23 variables involved in the development and maintenance of raised fields beyond agricultural  
24 practices, considering the hydrologic and climatic conditions in different geographical  
25 locations.

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27 In the floodplain wetlands, the flood pulse is between the main environmental driving forces  
28 that link nutrient fluxes to primary production, hydrobiological resources and economic  
29 exchange (Junk et al., 1989; Montoya et al., 2011). The changes in frequency and intensity  
30 of flood pulses owing to climate variability and anthropic interventions have shaped the  
31 landscape and nutrient dynamics (Junk et al., 1989; Macklin and Lewin, 2015). In the Gulf

1 of Urabá, the performance of hydrological resources depends on the regional water cycle  
2 which includes marine and hydrogeological processes as well as the local geomorphology  
3 influenced by anthropic mounds, channels, platforms and ridges of prehispanic origin  
4 (Denevan, 2006; Lombardo et al., 2011; Palomino-Ángel et al., 2019; Poveda and Mesa,  
5 1999; Rodrigues et al., 2014). This is particularly relevant because it affects the biocoenosis  
6 and nutrition loads for fishing, hunting, gathering or cultivation (Merten et al., 2020; Muller,  
7 1995; Parolin and Wittmann, 2010; Vasey et al., 1984), leading to cultural changes and  
8 particular landscape and territorial patterns.

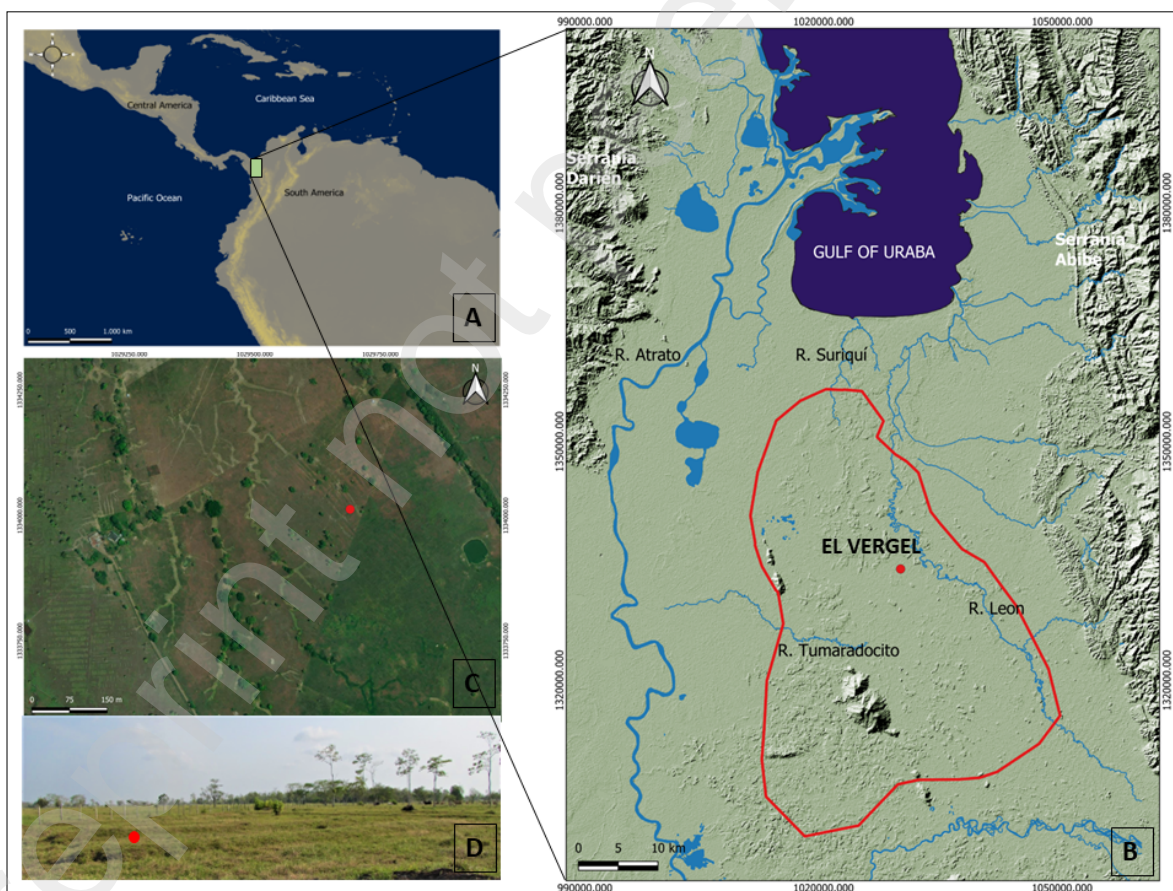
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10 The archaeology of the Gulf of Urabá in this regard, show prehispanic populations in the  
11 transition zone between Central and South America who shared phylogenetics and cultural  
12 features (Capodiferro et al., 2021; Martín et al., 2015; Mendizábal et al., 2021; Piazzini,  
13 2020), but material culture in the León river on the southwest of Urabá is clearly different  
14 from that of neighboring populations to the northeastern (Alzate, 2015; Arcila, 1955; GIAP,  
15 1980; Posada et al., 2019; Santos, 1989). Similarly, the wetlands are restricted to the southern  
16 part of the gulf where precipitation are higher and topography favors the permanent flooding  
17 by river overflow and water table fluctuations. This points out the relevance of the space-  
18 time framework for the Urabá raised fields to elucidating distribution, performance and origin  
19 of this technology and the cultural practices for water use. So this study explores the  
20 paleoenvironmental conditions in El Vergel archaeological site, focusing on the  
21 hydroclimatic complexity of Urabá and its role in the cultural practices, as the first approach  
22 to the human occupation history of the León river wetlands and its lessons for contemporary  
23 water management.

## 24 25 **2. Geographical settings**

26 El Vergel archaeological site lies in the middle reaches of the León river, 4 km West from  
27 the river stream (Fig. 1). It is a massive wetland filled with ditches, platforms, ridges and  
28 mounds that are enclosed by the mountain ranges of Darién to the West and of Abibe to the  
29 East. The wetland links the watersheds of the rivers Atrato, León, Tumaradocito and Suriquí  
30 within a relatively homogeneous swampy lowland. The geology consists of basal  
31 sedimentary rocks from Paleogene-Neogene age comprising lithic to sublithic sandstones

1 interspersed with mud and limestone. Quaternary alluvial sediments rest on the sedimentary  
2 layers and together form a multilayer aquifer with different levels of porosity and  
3 permeability (IDEAM, 2013). The unconsolidated sediments derived mainly from the Abibe  
4 mountain range, consist of sands and silty-clays of silicate, calcareous and oxide minerals.  
5 In addition, igneous sediments from basalts, breccias, agglomerates, granites and andesitic  
6 porphyries are transported by the Atrato river from the Darién mountain range and some  
7 isolated hills. The sediment load to the León river basin from both mountain ranges totals  
8 with 810,000 ton/a (Guevara et al., 2015; Guzmán and Ceballos, 2001). Two granitic hills  
9 alone stand out inside this massive floodplain fringed by the raised fields, El Cuchillo and  
10 Lomas Aisladas.

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1 **Fig. 1.** Map of the raised fields distribution (red polygon) around El Vergel archaeological  
2 excavation (red dot); (A) Continental context; (B) Regional location; (C) Aerial image and  
3 (D) *in situ* photograph.

4 The wetland soils in the middle and lower reaches of the León river are periodically flooded  
5 during high precipitation periods, promptly recharging the two regional aquifers found  
6 underneath and increasing the runoff (Betancur-Vargas et al., 2017; Ríos and Martínez, 1995;  
7 Villegas et al., 2013). The average rate for annual rainfall varies between 2,000 mm to the  
8 North and 4,000 mm to the South, depending on the position of the Intertropical Convergence  
9 Zone (ITCZ). The precipitation pattern is unimodal with a dry season between January-  
10 March when the ITCZ is at its southernmost position, and a rainy season between April-  
11 December when the ITCZ is above the study area. The minimum rainfall rate amounts to 70  
12 mm/month and coincides with Caribbean trade winds in the dry season. The maximum  
13 rainfall rate surpasses 200 mm/month in the rainy season (Betancur-Vargas et al., 2017;  
14 Guzmán and Ceballos, 2001). According to the climatic and limnometric station located in  
15 Barranquillita, close to El Vergel archeological site, the average multiannual flow of the León  
16 river is 70.7 m<sup>3</sup>/s, but increases downstream due to other tributaries.

### 17 **3. Methods**

18 Fifteen control pits were surveyed by digging in the best preserved zone of El Vergel  
19 archaeological site along a West-East transect. Six out of 15 diggings shared artifact and  
20 charcoal evidence embedded in similar soil profile, so we chose a ridge and a channel around  
21 it for excavation. Stratigraphic integrity and earthwork pattern were considered while  
22 recovering soil, artifacts and charcoal samples during the excavation. Since the channel  
23 represents a local scale micro-depositional basin, the archaeological materials were best  
24 stratified in the channel sediments. The soil horizon nomenclature followed the parameters  
25 of the USDA soil taxonomy (Soil Survey Staff –SSS, 2017, 1999), while the stratigraphy  
26 adopted the identification and description criteria established in Goldberg and McPhail  
27 (2006), and Sloss and Krumbein (1969).

28

#### 29 **3.1. Palynology and diatom sampling, preparation and identification**

1 About 2 cm<sup>3</sup> sediment samples were collected at different depths according to stratigraphic  
2 distribution, organic matter and soil development. All samples were mechanically and  
3 chemically dispersed with sodium pyrophosphate (Na<sub>4</sub>P<sub>2</sub>O<sub>7</sub>) and then oxidized with  
4 hydrogen peroxide (H<sub>2</sub>O<sub>2</sub>). The sampling, preparation, identification and quantification of  
5 diatoms were performed after González (2022) according with Segecin et al. (1999). Only  
6 the pollen samples were subjected to acid hydrolysis by adding hydrofluoric (HF) and  
7 hydrochloric (HCL) acids (Faegri and Iversen, 1975). The material was warm-dried onto  
8 glass slides and mounted in Entellan® resin. Observations, measurements and  
9 photomicrographs were taken at 40x and 100x (Olympus CX41 microscope with AxioCam  
10 ERs 5s camera). Some photomicrographs were taken using a Scanning Electron Microscope.  
11 Owing to the differential preservation of diatoms down core, the taphonomic processes were  
12 analyzed qualitatively by fragmentation, dissolution and lamination of valves. The whole  
13 results, including artifacts and elemental geochemistry, were plotted with Grapher 13.3.754.

14

### 15 **3.2. Soil chemistry and micromorphology**

16 The physical and chemical properties of the soil sampled were compared with natural soils  
17 of the same physiographic unit in the León river basin reported in IGAC (Instituto Geográfico  
18 Agustín Codazzi - IGAC, 2007). The content of Fe, Ti and Al in alluvial sediments was  
19 determined by Inductively Coupled Plasma - Optical Emission Spectrometry (ICP-OES)  
20 after acid digestion. The organic matter content was estimated using the Walkley-Black  
21 method. Effective cation exchange capacity (ECEC) and base cations in soil (Ca, K, Na y  
22 Mg) were measured by absorption and emission spectrometry after extraction with  
23 ammonium acetate (NH<sub>4</sub>CH<sub>3</sub>CO<sub>2</sub>) 1M at pH 7. Total phosphorus was determined using  
24 potassium nitrate (KNO<sub>3</sub>)/sodium nitrate (NaNO<sub>3</sub>) and a spectrophotometer. Finally, soil  
25 micromorphology was evaluated following Bullock et al. (1985), with emphasis in  
26 archaeological features described in Courty et al. (1989), Stoops and Marcelino (2010) and  
27 Stoops and Nicosia (2020).

28

### 29 **3.3. Radiocarbon dates**

30 The <sup>14</sup>C dates were obtained by accelerator mass spectrometry (AMS) from charcoal samples  
31 collected in stratigraphic units without post-depositional disturbance. Calibrated dates were

1 obtained with the Calib 8.2 software using IntCal20 calibration curve and a  $2\sigma$  probability  
2 interval (Reimer et al., 2020; Stuiver and Reimer, 1993).

3

### 4 **3.4. Statistical analysis**

5 Principal component analysis (PCA) was performed in order to explain the variability of the  
6 geochemical, biological and artifact findings. Significant correlations between these  
7 variables were obtained with the Pearson test. All statistical analyses were carried out using  
8 IBM SPSS version 25 and Statgraphics 19 Centurion.

9

## 10 **4. Results**

### 11 **4.1. Archeological excavation, stratigraphy and chronology**

12 A 4 m trench 1,7 m deep was dug between the selected channel and the ridge, which displayed  
13 five stratigraphic units in a moderately developed soil (Fig. 2). The borders between units  
14 were diffuse due to weathering, with the exception of a clear discontinuity between  
15 stratigraphic units II, III and IV (Table 1). The distribution of ceramic fragments along the  
16 sequence was multimodal, with increased density after unit I where we record a subtle  
17 typological change. Unit I share similar chronology with the most lower level of Toribio  
18 archaeological site dated to about  $^{14}\text{C}$  1570 $\pm$ 53 yr BP (Posada, 2022). Pedological evolution  
19 was incipient in all units with absence of buried soil or paleosols in the whole sequence.

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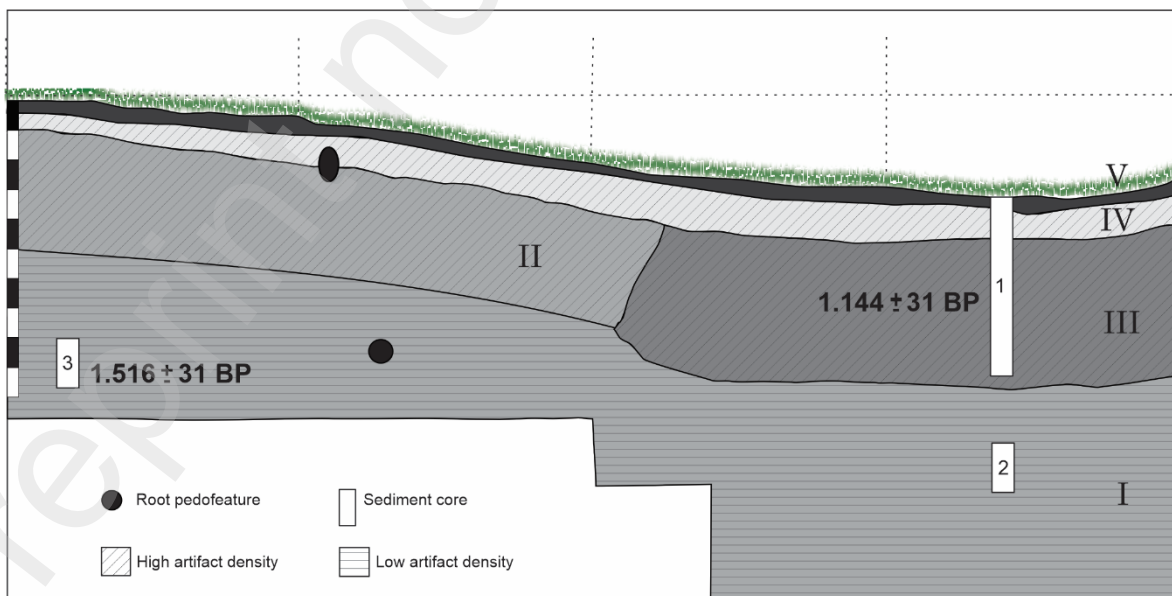
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1 **Fig. 2.** Stratigraphic sequence in El Vergel excavation

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Stratigraphic unit	Description
I (140-60 cm)	Clayey-silt sediment layer, densely mottled with abundant Fe/Mn oxide nodules forming a Cgu soil horizon bearing charcoal and artifacts. Its upper limit to the ridge was flat and diffuse but wavy and abrupt to the channel. Radiocarbon age at the top layer of the ridge was 531-609 CE ( $^{14}\text{C}$ 1516 $\pm$ 31 yr BP; AA115152).
II (60-15 cm)	Loamy-clayed sediment layer occurred in the ridge only and displayed increased artifact density. The color was yellowed-brown with dispersed Fe/Mn oxide nodules forming a Cgu soil horizon. Root pedofeatures and charcoal were frequent. The upper limit was wavy and gradual but laterally discordant with unit III.
III (60-20 cm)	Channel filling clay-loam sediments, ceramic artifacts and charcoal with less pedological development. Gleyed mottle morphology with Fe/Mn oxide nodules, massive structure and root pedofeatures. The upper limit was discordant with the upper unit IV. Radiocarbon age at the middle of the unit was 868-991 CE ( $^{14}\text{C}$ 1144 $\pm$ 31 yr BP; AA115150).
IV (20-5 cm)	This layer covered the El Vergel archaeological site throughout. It was dominated by loamy-sand sediments, pale gray color with some mottles, contained ceramic artifacts and few charcoal with no Fe/Mn oxide nodules. The upper limit was flat and abrupt, suggesting a discontinuity with the next stratum. Radiocarbon age was post bomb (1962-1977 CE), which along with the fluvial origin of sediments may suggest reworked materials.
V (5-0 cm)	Thin layer of loamy soil, brown-gray color with some mottles, strongly compacted structure and mid to high content of organic matter forming an Ap horizon. Both ceramic artifacts and charcoal were absent.

3 **Table 1.** Pedostratigraphical description.

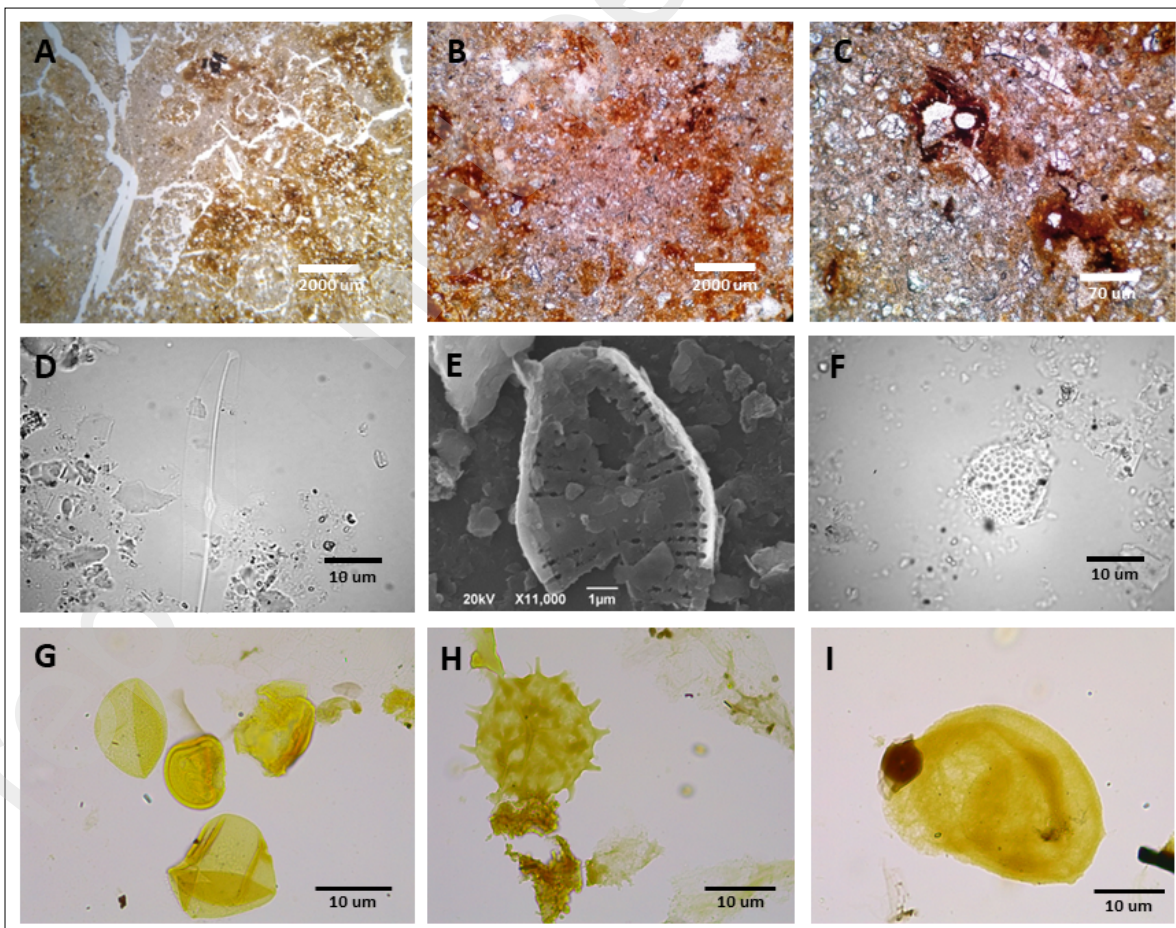
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5 **4.2. Site formation processes**

6 Hydromorphism was the main pedogenic process of the archaeological site. This is consistent  
 7 with the high presence of Fe/Mn oxide nodules, ferruginous hypo-coatings around root voids  
 8 and gleyed zones of the soil profile (Fig. 3). These features are produced *in situ* by the  
 9 fluctuation of groundwater levels that lead to degradation of palynomorphs, organic and  
 10 carbonate debris as those associated with faunal bones and shells. The steep reduction in  
 11 diatom frequency under stratigraphic unit IV was probably due to dissolution of fragile  
 12 frustules under long term waterlogging variations, opal hydration and mechanical  
 13 fragmentation during transport and deposition (Barker, 1992; Flower, 1993; Reed, 1996;  
 14 Sierra-Arango et al., 2014; Warnock et al., 2007). Because the ascension of alkaline  
 15 groundwater and marine intrusions through the riverbed (Jiménez and Campillo, 2020;

1 Villegas et al., 2013), the chemistry of the channel soil stabilized at alkaline pH levels and  
2 were clearly depicted by the increased downcore content of  $\text{Ca}^{2+}$  and  $\text{Mg}^{2+}$ , measured in the  
3 stratigraphic units I, II and III, where  $\text{Mg}^{2+}$  saturation reached high levels (37-44 %; Table  
4 2). Such pH levels increase silica dissolution of bioclasts despite being more resistant than  
5 carbonated (Fritz et al., 1999; Jorgensen, 1955; Ramírez et al., 2007).

6  
7 Another relevant processes were the flood pulse and the water flow inside the channel. These  
8 featured diverse minerals, mild chemical weathering, homogeneous soil morphology, no  
9 buried A horizon and low levels of sedimentary organic matter (Table 2; Fig. 2). This concurs  
10 well with frequent but moderate water fluxes into the channel and ridge. Both processes cause  
11 further fractionation of diatoms, pollen oxidation, variation in trace elements levels and loss  
12 of suspended clay (Sierra-Arango, 2013). The preservation of the archaeological record is  
13 likewise affected by mechanical bioturbation observed in roots, microfaunal remains and  
14 fabrics (Fig. 3A). Undoubtedly, the exceptional event registered in the stratigraphic unit IV  
15 would have caused reworking and mixing ages of the layers.



1

2

3 **Fig. 3.** Palynological and soil micromorphological features. (A) Loose continuous  
4 excremental infillings in stratigraphic unit I; (B) Weakly developed pedality in stratigraphic  
5 unit IV; (C) Fe oxide hypocoatings in a yellowish brown groundmass with dense angular  
6 quartz crystals in stratigraphic unit III; (D) *Gyrosigma scalproides* diatom; (E) *Luticola*  
7 *acidoclinata* diatom; (F) *Aulacoseira* spp diatom; (G) Poaceae pollen grains; (H)  
8 Convulvulaceae pollen grain; (I) Podocarpaceae pollen grain.

Channel data																
Stratigraphic Unit	pH	% Organic carbon	% Organic nitrogen	C/N Ratio	Ca (meq/100 g)	K (meq/10 g)	Mg (meq/10 g)	Na (meq/10 g)	ECE C (meq/100 g)	Total P (mg/K g)	Cation Saturation (%)	Ca/Mg Ratio	K/Mg Ratio	Clay (%)	Silt (%)	Sand (%)
V	5.18	5.04	0.59	11.59	4.68	0.46	2.9	0.16	9.06	251.2	35.39	1.60	6.34	42.2	45.4	12.4
IV	6.19	0.81	0.10	11.59	6.49	0.06	3.76	0.14	10.45	91.57	59.98	1.72	62.66	24.1	45.2	30.7
III	7.37	0.29	0.03	11.61	8.39	0.09	7.68	0.19	16.35	88.51	86.42	1.09	85.33	26.0	41.1	32.9
III	7.38	0.23	0.03	11.60	12.29	0.18	8.67	0.15	21.29	87.92	91.18	1.41	48.16	13.2	45.5	41.3
I	7.45	0.38	0.05	11.60	9.73	0.12	8.64	0.15	18.64	127.05	80.98	1.12	72.00	19.6	49.5	30.9

Ridge data																
Stratigraphic Unit	pH	% Organic carbon	% Organic nitrogen	C/N Ratio	Ca (meq/100 g)	K (meq/10 g)	Mg (meq/10 g)	Na (meq/10 g)	ECE C (meq/100 g)	Total P (mg/kg)	% Cation Saturation	Ca/Mg Ratio	K/Mg Ratio	Clay	Silt	Sand
V	5.68	3.50	0.41	11.59	6.71	0.23	6.32	0.35	13.61	130.40	62.26	1.06	27.47	48.5	37.1	14.4
IV	7.25	0.24	0.03	11.58	7.23	0.05	7.11	0.31	14.70	76.22	80.58	1.01	142.20	30.2	43.1	26.7
II	7.44	0.32	0.04	11.59	9.56	0.06	8.38	0.20	18.20	88.19	91.72	1.14	139.66	30.1	45.3	24.6
II	7.27	0.30	0.04	11.61	12.72	0.09	8.59	0.13	-	78.03	Saturated	1.48	95.44	32.2	43.2	24.6

I	7.35	0.39	0.05	11.59	13.30	0.11	8.71	0.12	-	75.30	Saturated	1.52	79.18	27.8	45.4	26.8
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**Table 2.** Physical and chemical properties of soil in El Vergel pedostratigraphic sequence.

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### 1 4.3. Palynomorphs and biosiliceous components

2 The pollen sequence showed ecological patterns and successional changes, but no evidence  
3 of agricultural crops. In the stratigraphic unit I, 82% of the pollen recorded at 180 and 100  
4 cm corresponded to tropical rainforest taxa highlighted by *Forsteronia*, *Juglans* sp.,  
5 *Desmodium* sp., *Alnus* sp. and Apocynaceae (Fig. 3). While the savannah pollen got 4.5%  
6 represented by Poaceae, and finally the swamp taxa 13.5% with Cyperaceae. A low diatom  
7 abundance with alternating dominance was recorded. *Aulacoseira* sp. dominated between 25-  
8 100 cm and *Gyrosigma* cf. *scalproides* below 100 cm. Some resistant structures from centric  
9 diatoms, *Actinella*, fragments of *Eunotia* cf. *monodon*, *Pinnularia* sp., *Stauroneis* sp. as well  
10 as some sponge spicules were also found in the unit I.

11  
12 The stratigraphic unit III, was dominated by rainforest group and represented by *Carya* sp.  
13 Type, *Tapirira* sp. Type, *Podocarpus* sp. and *Forsteronia* sp. Cyperaceae increase the swamp  
14 taxa (24%) followed by savannah vegetation (17.1%) with Poaceae. Although Fabaceae and  
15 Anacardiaceae were the dominant rainforest taxa, their count decreased atop the core whilst  
16 Poaceae and Asteraceae increased. Further analysis identified some *Luticola incana* and  
17 *Pinnularia borealis* from the algal group atop the stratigraphic unit III (Fig. 3).

18  
19 In the stratigraphic unit IV, open vegetation increased with Poaceae and *Ambrosia* sp.  
20 representing 38.3%, whereas Cyperaceae, Onagraceae and *Typha* type grains reached 31.3%,  
21 suggesting a transitional environment dominated by savannah and swamp vegetation with  
22 few rainforest taxa such as *Palicourea* sp., *Juglans* sp., *Desmodium* sp. and Arecaceae. The  
23 diatom assemblage at 17 cm deep was represented by *Luticola acidoclinata*, *Pinnularia*  
24 *borealis*, *Eunotia* and *Aulacoseira* sp., whereas the diatom *Brachysira* sp. only occurred in  
25 the record at 15 cm.

26  
27 The last stratigraphic unit registered high pollen frequencies from both savannah and swamp  
28 taxa. Whereas savannah comprised 44.5% with Poaceae, *Ambrosia* sp. and Asteraceae,  
29 swamp included Cyperaceae, Onagraceae and *Thypha* type grains with 32%. Nevertheless, the  
30 rainforest ecosystem recovered slightly at 3 cm by increasing its diversity with Bromeliaceae,  
31 *Alchornea* sp., *Acalypha*, *Desmodium* sp., *Trema* sp., *Hedyosmum* sp., *Pseudobombax* sp.

1 and *Symphonia*. The richness of well preserved diatom taxa was high, being the genus  
2 *Pinnularia*, *Luticola*, *Frustulia*, *Nitzschia*, *Hantzschia*, *Eunotia*, *Sellaphora*, *Encyonopsis*,  
3 *Stauroneis*, *Navicula*, *Gomphonema*, *Placoneis* and *Achnantes* the most conspicuous  
4 (González, 2022).

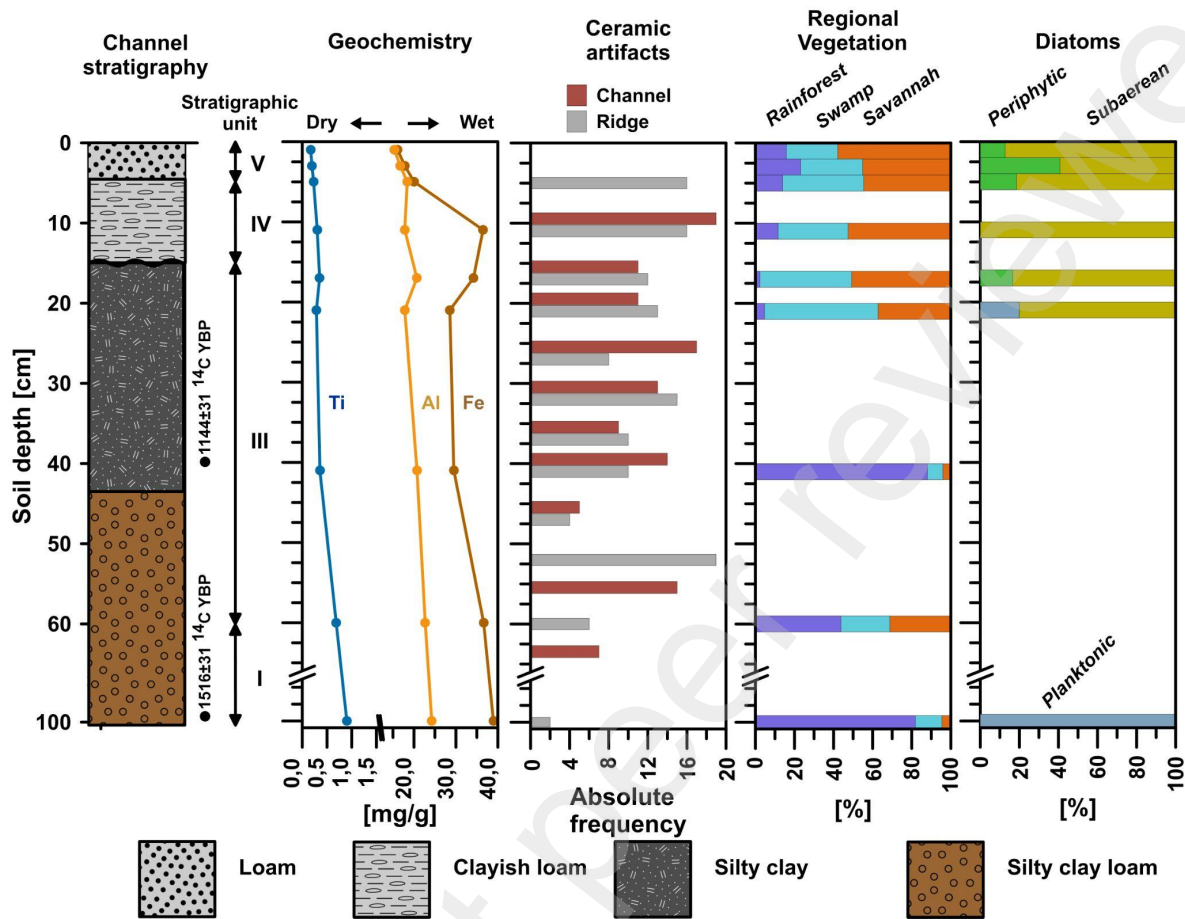
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#### 6 **4.4. Geochemistry and soil micromorphology**

7 The sedimentary content of Fe, Al and Ti increased down core with one peak resolved  
8 between 10-20 cm depth (Fig. 4). The levels of Fe, Al and Ti were higher at the bottom core  
9 in unit I than in the rest of the sequence, indicating increased runoff. The levels of Fe and Al  
10 were higher than that of Ti. The micromorphology exhibited diverse biostructures and  
11 inorganic sediments slightly altered, well developed porosity and several impregnative  
12 Fe/Mn oxide nodules, some of them cut forward by pedofaunal activity (Fig. 3A). The  
13 organic carbon was very low although total phosphorus (P) reached 127 mg/kg in the deepest  
14 stratum. The levels of Fe, Al and Ti decreased gradually from 40 cm in stratigraphic unit III  
15 to 20 cm depth. The period of greatest artifact deposition within the channel load  
16 corresponded with 40 cm depth ( $^{14}\text{C}$  1144±31 yr BP). The ECEC (Effective Cation Exchange  
17 Capacity), base saturation and pH levels were higher in the deeper units than those reported  
18 for natural and recent soils (Instituto Geográfico Agustín Codazzi - IGAC, 2007), but may  
19 be accentuated by groundwater salts (Villegas et al., 2013).

20

21 The features of the stratigraphic unit IV suggested a sudden flood event between 20-5 cm,  
22 when the levels of the three elements peaked. These levels of Fe and Al almost match those  
23 found in the stratigraphic unit I before anthropic occupation. However, these should be seen  
24 with caution since the apparent increase in humidity might be partially overestimated by  
25 redox conditions. In fact, the soil micromorphology showed diverse redoximorphic features,  
26 including intrusive and impregnative pedofeatures. Soil microstructures displayed less  
27 porosity and pedality than the stratigraphic unit I, with frequent root remains, coarse  
28 minerals, nucleic organic nodules, and occasional occurrence of micromollusks (Fig. 3B).  
29 The peaks of the three metals dropped suddenly between 12-1 cm to levels indicating  
30 desiccation of the León River wetlands in modern times. Accordingly, the pH, ECEC and  
31 base saturation decreased as opposed to the organic matter and total phosphorus (Table 2).



2

3

4 **Fig. 4.** Vertical distribution of multiproxy data from El Vergel channel core.

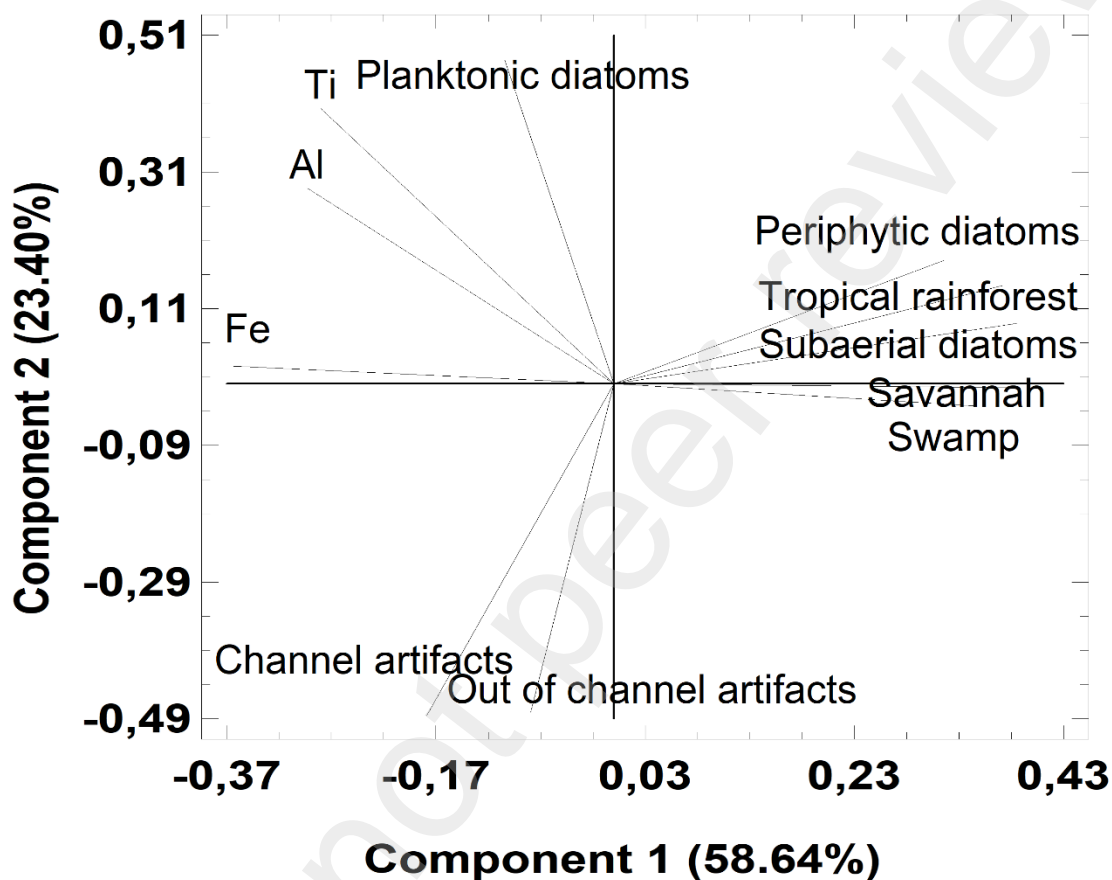
5

6 **4.5. Data correlation and paleoecological reconstruction**

7 The vertical and statistical correlation of data show the reliance of pollen, diatoms and  
 8 geochemistry for hydroclimatic interpretation in spite of Fe oxidation. The variance  
 9 explained by two principal components is 82.04% (Fig. 5). The component 1 explained  
 10 58.64% of the variance and was mostly weighed by Fe, pollen and subaerial and periphytic  
 11 diatoms. The component 2 was represented by Al, Ti, planktonic diatoms and ceramic  
 12 artifacts. The contents of Fe, Al and Ti correlated significantly ( $p < 0.05$ ) among them with  
 13 Pearson correlation coefficients ranging 0.76-0.92. The Fe reverse correlated ( $p < 0.05$ ) with  
 14 subaerial and periphytic diatoms as well as pollen, with Pearson correlation coefficients  
 15 between 0.67-0.88. Neither Al or Ti correlated significantly with biological data. Pollen of



1 the three assemblages correlated highly ( $p < 0.01$ ) among them and with subaerial diatoms  
 2 (Pearson  $r = 0.85-0.97$ ). The periphytic diatoms correlated ( $p < 0.05$ ) with subaerial diatoms as  
 3 well as pollen from tropical rainforest and savannah (Pearson  $r = 0.67-0.89$ ). No significant  
 4 correlation was found between ceramic density and chemical or biological data.



23 **Fig. 5.** Principal Component Analysis for pollen assemblages, diatoms, artifacts and  
 24 geochemistry.

26 The sediment load to the floodplain arises from weathering of the densely forested mountains  
 27 of Serranía de Abibe, therefore the content Fe, Al and Ti in channel sediments can be used  
 28 for elucidating the frequency of fluvial events by hydroclimatic forcing. Indeed, high  
 29 sedimentary content of Fe, Al and Ti was found in the channel along with high pollen density  
 30 from tropical rainforest and low density for both human pottery and planktonic diatoms (Fig.  
 31 5).

1 The human occupation began at 531-609 CE before channel development, when the record  
2 of *Forsteronia* sp. woody creepers in the palynomorphs indicated prevailing humid  
3 conditions with a long term decreasing trend towards present. According to artifact  
4 deposition, this first human incursion was not permanent but stationary. The prevailing  
5 humid conditions depicted by the palynomorphs concur well with high levels of Fe, Al and  
6 Ti derived by high sedimentation rates and runoff. The amount of diatom valves in the  
7 channel soil was low with sparing occurrence of *Aulacoseira* sp. and *Gyrosigma* cf.  
8 *scalproides* (Fig. 4). The raised fields were built before 868-991 CE, when human occupation  
9 became permanent, coinciding with decreased precipitation that caused diminished pollen  
10 accumulation rates although the rainforest remained yet the main local ecosystem. The low  
11 rates for accumulation of rainforest pollen continued until fading within the stratigraphic unit  
12 IV. The occurrence of subaerial diatoms indicated reduction in aquatic habitat as well.

13

14 The geochemistry suggested a light increase in runoff that was not confirmed by either pollen  
15 or diatom records during this period in the stratigraphic unit IV. This apparent contradiction  
16 needs to be interpreted in the light of newly increased erosion rates due to human earthworks  
17 and the mixing properties of this altered stratigraphical unit by an atypic flood event. Clearly,  
18 this event was out of control of the raised fields because it is widespread around the site on  
19 both the channel and ridge. The Fe peak atop this stratigraphic unit could be regarded as an  
20 effect of oxidation (Muñoz et al., 2017; Petersen et al., 2000) probably during plant  
21 diversification due to warming and dry weather. This peak was concomitant with a lull for  
22 Ti and Al, which could thus reliably be interpreted as decreased deposition of sediments.

23

24 In fact, diatoms were very scarce and forest regenerated slightly displaying other new taxa,  
25 but savannah vegetation persisted as long as artifacts deposition reached the highest count.  
26 The incipient soil development, the diversity of biological remains and redoximorphic  
27 features observed micromorphologically, along with chronostratigraphic data, suggest that  
28 the stratigraphic unit IV was caused by a local, strong and short term flood event during a  
29 dry and warm environment. The environmental conditions beyond the stratigraphic unit IV  
30 changed progressively to dryness and evinced people suddenly disappearing.

31

## 5. Discussion

The first humans arrived at El Vergel archaeological site around VI century CE, but the earthworks seem to have been developed for permanent use a few centuries later. Our paleoecological and chronostratigraphic data highlighted the prevalence of a wet environment in El Vergel during this first stage of occupation. Additionally, the data depicted a long term trend towards a dry period after construction of the raised fields around the IX century CE. It is consistent with palynological records from the western Colombian Andes (Muñoz et al., 2017; Rangel et al., 2005; Velásquez, 2005) and the Panamanian Darien (Bush and Colinvaux, 1994; Lachniet et al., 2004) during a fairly dry period between CE 1100 and 1200 year BP regarded as Medieval Warm Period (Acevedo et al., 2020; Berrío et al., 2000; Plazas et al., 1988; Van der Hammen et al., 2005; Velásquez, 1999). Hence the runoff reduction could not have the anthropic influence only.

Additional palynological data coming from the coastal plains is, however, challenging due to elution by large sediment loads collected in the lowlands (Lazala et al., 2010; Romero and Rangel, 2013). Other factors adding to this elution are the anthropogeomorphic disturbances and ENSO variability that result in increased precipitation, erosion and runoff (Haug et al., 2001; Martin et al., 1993; Muñoz et al., 2017). The sediment load arising from ENSO variability is particularly high in the southern Gulf of Urabá owing to massive precipitation caused by the convergence of the low level atmospheric jets Panama and Chocó beyond the seasonal migration of the intertropical convergence zone (ITCZ) (Poveda and Mesa, 1999; Rúa et al., 2015; Vélez-Agudelo and Aguirre-Ramírez, 2016). Besides climatic patterns, the León river basin is a groundwater dependent system which floods the wetland regardless of rainfall at short term (Arana, 2015; CORPOURABA, 2019), so regional hydrogeology plays an important role in the water storage and transfer to the ecosystem.

The whole data suggest a reduced, continued and slow deposition of sediments, without organic accumulation and mild mineral weathering associated with unstable geomorphic processes. Most of these processes were influenced by water table fluctuations and increased flow through artificial ditches. This led to incipient soil development and low fertility according to poor organic carbon and phosphorus contents, besides the potentially toxic

1 soluble salts enrichment in the soil. For this reason along with the lack of cultivars, clearing  
2 or plowing evidence, we cannot link the studied raised fields with agricultural practices.  
3 Conversely to other raised fields in South America with cultivation evidence (Berrío et al.,  
4 2001; Boixadera et al., 2019; Iriarte et al., 2010; Plazas et al., 1988; Rodrigues et al., 2020,  
5 2014), the El Vergel profile does not show either buried soils nor peat that could be related  
6 to organic rich/manure accumulation model (Lombardo et al., 2011).

7

8 The low content of sedimentary organic matter and coarse granulometry during channel  
9 operation, is probably because the water in the channels was not stagnant, but rather flowing,  
10 leading to suspension transport of low gravity sediments, accumulation of coarser ones and  
11 allowing mechanical diatom fragmentation. This is reinforced by the domination of mottled  
12 or redoximorphic features instead of gleyed morphology in the whole profile. Therefore,  
13 there is a clear probability that prehispanic communities in El Vergel site managed a climate  
14 transition towards a less humid environment by developing raised and ditched fields for  
15 reasons different from agricultural production (Lombardo et al., 2011; Morlon, 2006;  
16 Rodrigues et al., 2018).

17

18 The lack of agriculture should not be overestimated because of the restrictions of our survey  
19 and the differential performance of raised fields (Morlon, 2006; Renard et al., 2012).  
20 Nonetheless, the cultivated plants could be a marginal food resource regarding fishing, turtle  
21 farming, hunting and gathering as important strategies against local ecosystem services  
22 (Comptour et al., 2018; Erickson, 2000; Jaramillo and Jiménez, 2008; Márquez, 2008). By  
23 the way, Posada (2022) reports a domestic context in the Toribio archaeological site, near  
24 13.5 km from El Vergel, with hearth evidence and frequent remains of mollusks.

25

26 The use of ditched fields for managing flood pulses proposed by Lombardo et al. (2011),  
27 seems also plausible in El Vergel. However, there is evidence to support a wider hypothesis  
28 including the use of this technology for regional distribution of water during a period of  
29 hydric stress. Sourcing would be fluvial as well as subterranean from available shallow  
30 springs (Betancur-Vargas et al., 2020; Ríos and Martínez, 1995) and developed by wells and  
31 ditched fields, in the same way of Maya wells and reservoirs in seasonally dry regions

1 (Isendahl, 2011; Johnston, 2004). Water redistribution would be akin to that of natural flood  
2 pulses to mitigate progressive dryness, and could explain the pervasive occurrence of the  
3 raised fields with interconnected channels and no clear boundaries among the watersheds of  
4 the rivers León, Suriquí and Tumaradocito (Posada, 2022).

5  
6 This study is unable to provide a rationale for the practical disuse of these raised/ditched  
7 fields at the end of the prehispanic period. Nevertheless, the sudden event related to the  
8 occurrence of the stratigraphic unit IV must be taken into account, since it broke the historical  
9 sequence. Despite bearing potentially reworked materials at this stratum, accounting for the  
10 above mentioned strong depositional event together with both demographic increase and  
11 intense alteration of the landscape, constitute valuable clues for further research to understand  
12 the abandonment of the site before the Spanish conquest of Urabá

13  
14 After reviewing these ideas through multiproxy data and specific geographic conditions in  
15 the Gulf of Urabá, considerable insight has been gained in regard to the hydrological system  
16 management of prehispanic communities beyond agricultural practices. Thus, people  
17 profited on a wide range of hydrobiological resources taking advantage of the hydric and  
18 hydraulic predictability of the raised/ditched fields to reduce cultivation dependence and  
19 climatic change risks.

## 20 21 **6. Conclusions**

22 Our study at the El Vergel archaeological site shows that around 531- 609 yr CE, the first  
23 evidence of human occupation was recorded in the León River wetlands with no signs of  
24 earthworks stratigraphically associated. The occupation seems to be sporadic or seasonal  
25 according to the vertical variability of ceramic deposition. Prevailed humid climatic  
26 conditions and forest cover in the wetland at this stage.

27  
28 Geochemical, palynological, and diatom data reveal a progressive shift toward drier climatic  
29 conditions. At the same time, human occupation of the site intensifies, peaking around 868 –  
30 991 yr CE. The peak of the occupation occurs when channel and ridge have been developed  
31 earlier.

1

2 In spite of critical effects on the palynomorphs and biogenic silica bodies preservation, and  
3 the redoximorphic conditions, salt accumulation and fluvial transport processes, we found  
4 low levels of natural fertility in the soil by physical and chemical properties and no evidence  
5 of agriculture by macro and micromorphological features, so we cannot to establishing clear  
6 associations between raised fields and agricultural practices.

7

8 Instead, the presence of well-formed channels in the warming period, the spatial distribution  
9 of shallow springs from the aquifer and the report of aquatic faunal remains in the nearby  
10 archaeological site Toribio, suggest the development of raised fields for water use and  
11 management, probably for wetland conservation and aquaculture during drought period.  
12 These results highlight the role of raised fields for hydric control as the primary resource of  
13 the wetland ecosystem in Urabá, encouraging new lines on research for the Colombian-  
14 Isthmic archaeology, paleoecology and conservation.

15

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23

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